## May 2009 Atmospheric River Event in the Dronning Maud Land

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On the Antarctic continent, precipitation is the only significant source term in the ice sheet mass balance, counterbalancing ice discharge and basal melt (Krinner et al., 2007, Uotila et al., 2007). Precipitation is also particularly difficult to quantify in Antarctica - an extremely cold and dry environment, prone to high winds redistributing and sublimating snow.

Measurements by an automatic weather station installed in Dronning Maud Land at the northern foot of the Sør Rondane mountain range (at the Princess Elisabeth station, 71°057' S, 23°021' E, 1392m asl, 173 km inland) showed a particularly strong accumulation event on 19th May 2009 (Gorodetskaya et al., 2013). The mean sea level pressure map based on the European Center for Medium Weather Forecast reanalysis (ERA-Interim) data showed that this event was associated with a deep cyclone blocked on the east by a high pressure ridge. Figure 1 depicts this strong synoptic system as seen by the infrared satellite composite product, developed at the University of Wisconsin (Lazzara et al., 2011). The cyclone was long lived, and directed elevated amounts of moisture onto the continent along the 60°E meridian. Such systems have been observed in the Arctic (e.g. Inoue et al., 2010) but such analysis in the Antarctic is only just beginning.

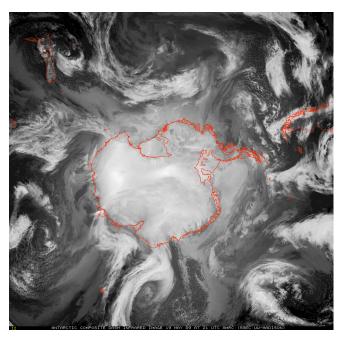


Figure 1. An infrared Antarctic composite image from May 19, 2009 at 21 UTC depicting a massive storm off Dronning Maud land (Courtesy of Dr. Matthew Lazzara, AMRC, University of Wisconsin).

SSM/I satellite data for May 19<sup>th</sup> 2009 reveal a narrow band of water vapor tracing meridionally across the Southern Ocean (Figure 2). Such a phenomenon is known as an "atmospheric river" in the Northern Hemisphere. Atmospheric rivers are relatively narrow

regions in the atmosphere that are responsible for significant midlatitude transport of water vapor (e.g. Ralph and Dettinger, 2011; Ralph et al., 2011). This phenomenon has been extensively studied in western US and is responsible for up to 50% of the annual precipitation in the western states. Although it has been hypothesized that atmospheric rivers are common in other parts of the world, the May 2009 case presents a first direct observation of an atmospheric river reaching the Antarctic continent.

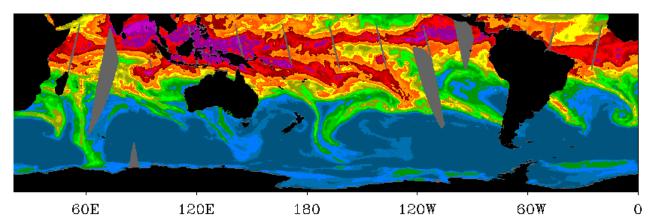


Figure 2. SSM/I Water Vapor (Wentz algorithm) for 19<sup>th</sup> May 2009, courtesy Gary Wick, NOAA. Note the narrow band of high water content around 60°E, stretching from the Indian Ocean mid-latitudes to the edge of the Antarctic continent.

Interestingly enough we became interested in May 2009 after performing the meridional moisture transport analysis (Tsukernik and Lynch, 2013) and noting a significant positive anomaly in total meridional moisture transport (Figure 3). Further calculations demonstrate that 18 separate synoptic events affected Dronning Maud Land in 2009, and that these accounted for 40% of the total annual accumulation at Princess Elisabeth AWS station that year (Gorodetskaya et al. 2013).

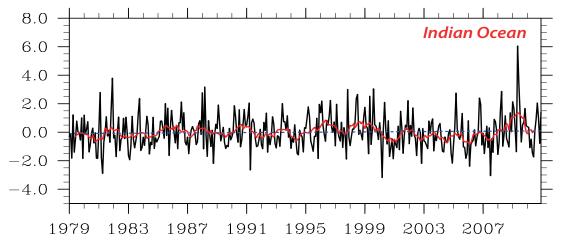


Figure 3. Transient eddy meridional moisture flux time series in the coastal zone of the Dronning Maud Land and surrounding area (Indian Ocean sector). Black line shows the monthly anomalies, red line shows interannual variability and the dashed blue line represents the long-term trend. Units are kg m-1 s -1.

In this study we performed a detailed analysis of the formation of the May 2009 atmospheric river. We employ the Weather Research and Forecasting (WRF) model to simulate the formation of the atmospheric river event. We also investigate the mechanisms important to the development of synoptic systems can be explored in an approximate form by using the quasi-geostrophic omega equation. In the quasi-geostrophic system, with no diabatic forcing, vertical velocity is forced solely by the divergence of the Q-vector (Hoskins et al. 1978). Departures from this regime can occur through contributions to the diabatic heating rate from surface sensible and latent heat fluxes and convection. We compute the relative contribution of these terms in the upper and lower levels (following the diagnostics developed by Lynch et al. 2003 and Bracegirdle and Gray 2008 for the Arctic cases) and discuss the implication of our results.

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